Global Earth Magnetic Field Modeling and Forecasting with Spherical Harmonics Decomposition

Panagiotis Tigas*  
OATML  
University of Oxford  
Oxford, UK

Téo Bloch*  
University of Reading  
Reading, UK

Vishal Upendran*  
IUCAA  
Pune, India

Banafsheh Ferdoushi*  
University of New Hampshire  
Durham, NH, USA

Yarin Gal  
OATML  
University of Oxford  
Oxford, UK

Siddha Ganju  
NVIDIA Corporation  
Santa Clara, CA, USA

Ryan M. McGranaghan  
ASTRA LLC  
Louisville, CO, USA

Mark C. M. Cheung  
Lockheed Martin  
Advanced Technology Center  
Palo Alto, CA, USA

Asti Bhatt  
SRI International  
Menlo Park, CA, USA

Abstract

Modeling and forecasting the solar wind-driven global magnetic field perturbations is an open challenge. Current approaches depend on simulations of computationally demanding models like the Magnetohydrodynamics (MHD) model or sampling spatially and temporally through sparse ground-based stations (SuperMAG). In this paper, we develop a Deep Learning model that forecasts in Spherical Harmonics space, replacing reliance on MHD models and providing global coverage at one-minute cadence, improving over the current state-of-the-art which relies on feature engineering. We evaluate the performance in SuperMAG dataset (improved by 14.53%) and MHD simulations (improved by 24.35%). Additionally, we evaluate the extrapolation performance of the spherical harmonics reconstruction based on sparse ground-based stations (SuperMAG), showing that spherical harmonics can reliably reconstruct the global magnetic field as evaluated on MHD simulation.

1 Introduction

The space environment around Earth (geospace) does not exist in a steady state. The dynamical changes in geospace is termed space weather. Space weather is primarily driven by interactions between the Sun’s output (in terms of solar radiation and the magnetized solar wind) with the Earth’s magnetosphere, thermosphere and ionosphere. The US National Oceanic and Atmospheric Administration continuously monitors the solar wind upstream from Earth with satellite observatories orbiting Lagrangian point L1 along the Sun-Earth line. The solar wind energy is transferred to the Earth’s magnetosphere via complex mechanisms [Dungey, 1961], leading to perturbations in the Earth’s magnetic field called geomagnetic storms.

*equal contribution

**We will release all the code and models used in this work post-publication.

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Solar wind parameters

Predictions
Sampled Predictions

Figure 1: Model architecture. Solar Wind data are summarized into an embedding vector using a Gated Recurrent Unit (GRU) which is then passed to a Multilayer Perceptron (MLP) to output the predicted spherical harmonics coefficients from where we can extrapolate values for specific locations.

Geomagnetic storms drive a spectrum of potentially catastrophic disruptions to our technologically-dependent society, among the most threatening being critical disturbances to the electrical grid in the form of geomagnetically induced currents (GICs). Due to their proprietary nature, publicly available GIC data are limited. However, a cohort study of insurance claims of electrical equipment provides evidence that space weather poses a continuous threat to electrical distribution grids via geomagnetic storms and GICs [Schrijver et al., 2014, Eastwood et al., 2018]. GICs also pose threats to oil pipelines, railways and telecommunication systems. In the case of extreme, but historically probable geomagnetic storms, the economic impact due to prolonged power outages can exceed billions of dollars per day [Oughton et al., 2017]. For this reason, there is urgency among public and industry stakeholders to improve monitoring and forecasting of space weather impacts like geomagnetic storms and GICs.

The geomagnetic field is continuously monitored, importantly with sparse spatial coverage, by a network of roughly 200 ground magnetometers [Gjerloev, 2012b] and by a constellation of 66 satellites collectively known as the Active Magnetosphere and Planetary Electrodynamics Response Experiment (AMPERE) project [AMPERE; Waters et al., 2020] in low Earth orbit. The spatially-sparse magnetometer measurements are typically synthesized into global indices (e.g. Dst, Kp, and AE; see appendix for more details) as measures of the geoeffectiveness of space weather perturbations. Like most indices, they are good as indicators but are too far removed from the underlying governing equations of the system. This makes interpretability and forecasting (whether by physics-based or physics-agnostic ML models) difficult.

GICs are driven by geoelectric fields through Ohm’s law, $\vec{J} = \sigma \vec{E}$ (the displacement current is negligible), where $\vec{J}$ is the current, $\sigma$ the conductivity tensor, and $\vec{E}$ is the electric field. The geoelectric field is related to the ground magnetic field perturbations through the Faraday’s law of induction $-\nabla \times \vec{E} = \frac{\partial \vec{B}}{\partial t}$, where, $\frac{\partial \vec{B}}{\partial t}$ is the rate of change of magnetic field on the Earth’s surface and can be derived from the time derivative of magnetic perturbation based on ground Earth magnetometers.

Our contributions are:

1. Using simulation data from a physics-based (MHD) model to validate a compressed sensing technique to recover global maps (in spherical harmonic basis) of the geomagnetic perturbation from sparse measurements. This improves the temporal cadence of such maps by >10x.

2. Developing a Deep Learning model that operates in the spherical harmonics space, allowing for global modeling of the magnetic field disturbances, combined with powerful non-linear autoregressive models (RNN, 1D-CNN) to capture the influence of solar wind data.

2 Datasets

We used data describing Earth’s magnetosphere and solar wind properties between 1 Jan 2013 and 31 Dec 2013. We used OMNI dataset for solar wind data, which captures the interplanetary magnetic field components (IMF), velocity and temperature of the solar wind as well as the clock.

2
angle. For Earth’s magnetic field, we used SuperMAG dataset [Gjerloev, 2012a] which consists of geomagnetic perturbations measurements from ground earth stations located at various places around the globe.

Finally, we used a simulation-derived dataset based on simulations conducted with Open Geospace General Circulation Model (Open GGCM) [Raeder et al., 2001], a magnetohydrodynamics (MHD) model. Lacking a global and spatially-complete ground-based magnetometer array we must rely on the first principles MHD model to constitute a global ground truth dataset. Furthermore, this allows us to validate our approach on finer resolution than one supported by SuperMAG and AMPERE dataset. However, note that the MHD model is not a substitute for actual perturbation measurements, for many short-scale phenomena are missed by such models [Raeder et al., 2001].

3 Methodology

In order to properly forecast GICs, we need to be able to model the relationship between the solar wind parameters and interplanetary magnetic fields (IMF) with the Earth system. Due to the scarcity of GIC data, the comparatively abundant magnetometer data and the relationship between the two, a logical step is to create models to predict $\frac{\partial \vec{B}}{\partial t}$. Such models provide complete spatial coverage and permit the study of the nature of the connection between the solar wind and the Earth.

3.1 Spherical Harmonics Decomposition

Any scalar field over the unit sphere can be expressed as $f(\theta, \phi) = \sum_{n=0}^{\infty} \sum_{m=-n}^{n} a_{nm} Y_{nm}(\theta, \phi)$, where

$$Y_{nm}(\theta, \phi) := \frac{2n + 1}{4\pi} \frac{(n - m)!}{(n + m)!} e^{im\phi} P_{nm}(\cos(\phi)),$$

and $P_{nm}(\cos(\phi))$ are the associated Legendre polynomials. These functions $Y_{nm}(\theta, \phi)$ are solutions to Laplace equation in a 3-D spherically symmetric coordinate system. If the sum is truncated at maximum harmonic degree $N$, $f(\theta, \phi)$ is approximated as $\tilde{f}(\theta, \phi) = \sum_{n=0}^{N} \sum_{m=-n}^{n} a_{nm} Y_{nm}(\theta, \phi)$.

Defining $i = n^2 + n + m$, the double sum can be written as $\tilde{f}(\theta, \phi) = \sum_{i=(N+1)^2-1}^{(N+1)^2-1} a_i Y_{i}(\theta, \phi)$. If the 2D fields over $\theta, \phi$ are unrolled as one-dimensional arrays, we have $\tilde{f} = \vec{B}\vec{a}$, where $\vec{a} = (a_i)$ is the tuple of spherical harmonic coefficients, and $\vec{B} = (\vec{b}_i)$ is the basis matrix wherein column vector $\vec{b}_i$ represents the basis function $Y_{nm}(\theta, \phi)$.

3.2 Experimental Design

3.2.1 Reconstruction

With SuperMAG data, we have spatially sparse samples $y$ which can be any component of the geomagnetic field, or its deviation from a reference field) measured at 200 stations. Note that only the Northern Hemisphere is considered in this analysis, for it has an extensive coverage (in contrast to the extremely sparse coverage in the Southern Hemisphere). The spherically symmetric reconstruction thus is $\tilde{f}(\theta, \phi)$, and this reconstruction as evaluated at the station locations is $\hat{y}$. Thus, we first construct a set of coefficients $\vec{a}$ from this sparse set of measurements, considered all the measurements across the globe at each time step. There are two constraints imposed for obtaining the set of coefficients:

1. The coefficient set must be sparse: This is to prevent power leakage and coupling across multiple modes.

$^3$ Solar wind measurement is performed by ACE and WIND satellites at the L1 Lagrange point of the Sun-Earth system.
2. As small \( N \) as possible: Since, in theory, an infinite number of modes can fit the observations, but a parsimonious model is required to limit overfitting. Furthermore, we would not want unnatural, localized artifacts due to high number of modes, thereby motivating a constraint on \( n \).

To mitigate constraint (1), we apply a Lasso regression technique [Tibshirani, 1996] with the spherical harmonic functions as the basis.

The Lasso regression comes with a regularization term \( \alpha \| \vec{a} \|_1 \), where \( \| \vec{a} \|_1 \) is the L1-norm of the coefficients and \( \alpha > 0 \) is a hyperparameter. This, and constraint (2) are mitigated by varying both \( \alpha \) and the maximum number of modes \( N \), and searching for a knee in a defined error metric, subject to the smallest maximum modes. The sweep parameters are detailed in Table 2, and two metrics are used – Maximum L1 error across all stations and time, and maximum R2 metric (coefficient of determination) across all time steps. The maximum L1 error tells us the worst performance across the dataset, and thus we seek the most acceptable worst possible performance. As shown in Fig. 3, the "knee of goodness" can be seen corresponding to \( \alpha = 0.1 \), and 20 maximum modes. These parameters are fixed in our analysis.

An example reconstruction of the data from 2017, considering stations above 40° is shown in Fig. 2. Here, we compare the north facing component of \( \partial d/\partial t \) from SuperMAG and its reconstruction.

Similarly, the comparison of spherical harmonic reconstruction for the MHD simulation is shown in Fig. 4. Note here, that the reconstruction is performed by sampling the MHD simulation at locations of SuperMAG stations alone.

### 3.2.2 Forecasting

Next, we construct a forecasting model which uses solar wind data (OMNI) to forecast the global magnetic field perturbation. For this experiment we use a similar setup as Weimer [2013] model. We feed 25 minutes of solar wind activity into a Gated Recurrent Network (GRU) to map the sequence into an embedding vector. We then feed the embedding into a Multilayer Perceptron (MLP) to output the spherical harmonics coefficients which model the global magnetic field perturbation; specifically we focus on north component of \( \partial d/\partial t \) as a proof of concept. In contrast to Weimer [2013] model, we use the whole sequence as input to a non-linear autoregressive model and we do not apply feature engineering to the OMNI data. Instead we use the raw features (see appendix for detailed list of features). The architecture can be seen in fig. 1.

We benchmark this work against the state-of-the-art empirical model by Wiemer et al. [Weimer, 2013]. To evaluate the performance we first compare on the validation set on SuperMAG but also we compare on simulations conducted with MHD model for two weeks worth of activity. The results are summarized in table 1.

### 4 Results And Discussion

<table>
<thead>
<tr>
<th>Model</th>
<th>SuperMAG (val) RMS (nT) ↓</th>
<th>MHD RMS (nT) ↓</th>
</tr>
</thead>
<tbody>
<tr>
<td>Ours.</td>
<td>24.23</td>
<td>27.02</td>
</tr>
<tr>
<td>Weimer [2013] model</td>
<td>28.35</td>
<td>35.72</td>
</tr>
</tbody>
</table>

With this work, we show and evaluate the reconstruction of the global magnetic perturbation field using spherical harmonics with LASSO regularization to promote sparsity in the coefficients. We show that it is possible to reconstruct from sparse measurements like the ones provided by SuperMAG stations,
Figure 4: Comparison of reconstruction with the MHD simulation. Left: MHD simulation; center: MHD simulation sampled at SuperMAG station locations; right: Reconstruction.

and we evaluated on MHD data to show that the reconstruction is similar to the global field modeled by MHD, suggesting that the results of such models can be compressed in spherical harmonics space by compressed sensing techniques. Additionally, we show that by using a Deep Neural Network to forecast in the spherical harmonics space, we can improve over existing state of the art (Weimer [2013] model) by 14.53% on SuperMAG and 24.35% on MHD dataset (summarized in table 1).

Broader Impact

Geomagnetic storms drive a spectrum of potentially catastrophic disruptions to our technologically-dependent society, among the most threatening being critical disturbances to the electrical grid in the form of geomagnetically induced currents (GICs). Due to their proprietary nature, publicly available GIC data are limited. However, a cohort study of insurance claims of electrical equipment provides evidence that space weather poses a continuous threat to electrical distribution grids via geomagnetic storms and GICs [Schrijver et al., 2014, Eastwood et al., 2018]. GICs also pose threats to oil pipelines, railways and telecommunication systems. In the case of extreme, but historically probable geomagnetic storms, the economic impact due to prolonged power outages can exceed billions of dollars per day [Oughton et al., 2017]. For this reason, there is urgency among public and industry stakeholders to improve monitoring and forecasting of space weather impacts like geomagnetic storms and GICs. With this work, we progress towards better and faster forecasting models which will help us shield our infrastructure from solar-wind related hazards.

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References


A Reconstruction experiment sweep parameter

<table>
<thead>
<tr>
<th>Sweep parameter</th>
<th>Description</th>
</tr>
</thead>
<tbody>
<tr>
<td>Date range</td>
<td>2015-06-23 to 2015-06-24</td>
</tr>
<tr>
<td>Cadence</td>
<td>10 min</td>
</tr>
<tr>
<td>Max modes sweep</td>
<td>[5,50], steps of 5</td>
</tr>
<tr>
<td>( \alpha ) sweep</td>
<td>( {2,1,0.1,1e^{-3},1e^{-5}} )</td>
</tr>
<tr>
<td>Max L1 metric</td>
<td>( \text{maxstations maxtime }</td>
</tr>
<tr>
<td>Max R2 score</td>
<td>( \text{maxtime } (1 - \sum_{\text{stations}}</td>
</tr>
</tbody>
</table>

Note that we have used a cadence of 10 min for quick computation of the metrics over the dataset — however, since the spherical harmonic generation is done at each time step, it can be performed at whatever cadence necessary.

B Solar Wind Data - OMNI Dataset

Here we describe the features we used from OMNI Dataset.

<table>
<thead>
<tr>
<th>Feature</th>
<th>Description</th>
</tr>
</thead>
<tbody>
<tr>
<td>( B_T ):</td>
<td>Magnetic field magnitude ( \sqrt{B_x^2 + B_y^2} )</td>
</tr>
<tr>
<td>( V_{SW} ):</td>
<td>Solar Wind velocity magnitude ( \sqrt{V_x^2 + V_y^2 + V_z^2} )</td>
</tr>
<tr>
<td>( T ):</td>
<td>Temperature of the solar wind</td>
</tr>
<tr>
<td>( \theta_c ):</td>
<td>Clock angle of the interplanetary magnetic field (IMF)</td>
</tr>
<tr>
<td>( F_{10.7} ):</td>
<td>F10.7 measures the noise level generated by the sun at a wavelength of 10.7 cm.</td>
</tr>
</tbody>
</table>

C Geoeffectiveness Indices

Dst: (Disturbance Storm Time Index) It is the measure of geomagnetic activity derived from near equator ground magnetic stations providing information about the strength of ring current.

Kp: Global geomagnetic index that is based on 3 hour measurements of mid-latitude ground magnetic stations around the world.

AE: (Auroral Electrojet Index) It is the measure of auroral activity determined based on ground magnetic stations around aurora zone.

D Weimer [2013] model

Weimer [2013] model baseline model is designed to forecast each of the three magnetic vector components. It uses solar wind data as input and it outputs the spherical harmonics coefficients, which then can be used to extract the forecasting values per CGM latitude and MLT pair.

A feature vector from the solar wind data is derived
\[
F = \begin{pmatrix}
1 & B_T & V_{SW} & t \\
\sqrt{F_{10.7}} & B_T \cos(\theta_c) & V_{SW} \cos(\theta_c) & t \cos(\theta_c) \\
V_{SW} \cos(\theta_c) & B_T \sin(\theta_c) & V_{SW} \sin(\theta_c) & t \sin(\theta_c) \\
\sqrt{F_{10.7}} \cos(\theta_c) & B_T \cos(2\theta_c) & V_{SW} \cos(2\theta_c) & B_T \sin(2\theta_c) \\
B_T \sin(\theta_c) & V_{SW} \sin(2\theta_c) & V_{SW} \sin(2\theta_c) & \end{pmatrix},
\]

where \(B_T\) represents the magnitude of the magnetic field, \(V_{SW}\) the solar wind velocity, \(\sqrt{F_{10.7}}\) the square root of the \(F_{10.7}\) feature (a measure of solar radiation), \(t\) the dipole axis angle in radians and \(\theta_c\) the clock angle.

Weimer [2013] model is trained on data obtained from SuperMAG stations for the year 2013 with the loss MSE = Avg((y - Ba)^2), where a are the spherical harmonics coefficients, computed as \(a_n^m = (g_n^m, h_n^m)\).

The coefficients are computed as \(g_n^m = G_n^m F\) and \(h_n^m = H_n^m F\), where \(G_n^m\) and \(H_n^m\) are the weights to learn.

The model is trained on 25 minute long average of solar wind data, with a lag of 20 minutes and as target the magnetic field perturbation of 5 minutes long averages, as measured by SuperMAG stations.

E Reproducibility details of forecasting experiments

The model architecture used for the forecasting experiment is presented below. The model was trained using Adam optimizer with learning rate \(lr=1e-04\) and Mean Squared Error as a loss.

```python
class GeoeffectiveNet(nn.Module):
    def __init__(self, 
                 past_omni_length, 
                 future_length, 
                 omni_features, 
                 supermag_features, 
                 nmax, 
                 targets):
        super(GeoeffectiveNet, self).__init__()

        self.omni_past_encoder = nn.GRU(25, 
                                         16, 
                                         num_layers=1, 
                                         bidirectional=False, 
                                         batch_first=True, 
                                         dropout=0.5)

        n_coeffs = 0
```
for n in range(nmax+1):
    for m in range(0, n+1):
        n_coeffs += 1
        n_coeffs *= 2

self.encoder_mlp = nn.Sequential(
    nn.Linear(16, 16),
    nn.ELU(inplace=True),
    nn.Dropout(p=0.5),
    nn.Linear(16, n_coeffs, bias=False)  # 882)

self.omni_features = omni_features
self.supermag_features = supermag_features

self.targets = targets
self.future_length = future_length

def forward(self,
past_omni,
past_supermag,
future_supermag,
dates,
future_dates,
**kargs):
    past_omni = NamedAccess(past_omni, self.omni_features)

    features = []
    # add the wiemer2013 features
    bt = (past_omni['by']**2 + past_omni['bz']**2)**.5
    v = (past_omni['vx']**2 + past_omni['vy']**2 + past_omni['vz']**2)**.5

    features.append(past_omni['bx'])
    features.append(past_omni['by'])
    features.append(past_omni['bz'])
    features.append(bt)
    features.append(v)
    features.append(past_omni['dipole'])
    features.append(torch.sqrt(past_omni['f107']))

    features.append(bt*torch.cos(past_omni['clock_angle']))
    features.append(v*torch.cos(past_omni['clock_angle']))
    features.append(past_omni['dipole']*torch.cos(2*past_omni['clock_angle']))
    features.append(torch.sqrt(past_omni['f107'])*torch.cos(2*past_omni['clock_angle']))

    features.append(bt*torch.sin(past_omni['clock_angle']))
    features.append(v*torch.sin(past_omni['clock_angle']))
    features.append(past_omni['dipole']*torch.sin(2*past_omni['clock_angle']))
    features.append(torch.sqrt(past_omni['f107'])*torch.sin(2*past_omni['clock_angle']))

    features.append(bt*torch.cos(2*past_omni['clock_angle']))
    features.append(v*torch.cos(2*past_omni['clock_angle']))
    features.append(past_omni['dipole']*torch.cos(2*2*past_omni['clock_angle']))
    features.append(torch.sqrt(past_omni['f107'])*torch.cos(2*2*past_omni['clock_angle']))

    features.append(bt*torch.sin(2*past_omni['clock_angle']))
    features.append(v*torch.sin(2*past_omni['clock_angle']))
    features.append(past_omni['dipole']*torch.sin(2*2*past_omni['clock_angle']))

    return features
features.append(torch.sqrt(past_omni['f107'])*torch.sin(2*past_omni['clock_angle']))
features.append(past_omni['clock_angle'])
features.append(past_omni['temperature'])

features = torch.stack(features, -1)
encoded = self.omni_past_encoder(features)[1][0]
coeffs = self.encoder_mlp(encoded)
predictions = torch.einsum('bij,bj->bi', future_supermag.squeeze(1), coeffs)

return coeffs, predictions, None

For the spherical harmonics decomposition we used scipy.special.sph_harm function.

def basis_matrix(nmax, theta, phi):
    from scipy.special import sph_harm
    assert(len(theta) == len(phi))
    basis = []
    for n in range(nmax+1):
        for m in range(-n,n+1):
            y_mn = sph_harm(m, n, theta, phi)
            basis.append(y_mn.real.ravel())
            basis.append(y_mn.imag.ravel())
    basis = np.array(basis)
    return basis.reshape(-1, theta.shape[0], theta.shape[1]).swapaxes(0, 1).swapaxes(2, 1)

F International Geomagnetic Reference Field

IGRF (International Geomagnetic Reference Field): Empirical measurements of the Earth magnetic field representing the main field without external sources.